Volcanic cloud retrievals from satellite using multispectral data in the TIR. Algorithm description, error assessment and perspectives

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- > Introduction
- > MODIS and SEVIRI multispectral satellite sensors
- > Ash and SO₂ detection
- > Ash and SO₂ retrievals
 - LUT approach
 - VPR approach
- > Volcanic Ice detection and retrievals
- Sensitivity analysis
- Validation
- Conclusions

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Volcanic Gas emissions

Particles

The term 'volcanic ash' refers to particles having dimensions varying between a few mm and less than 1 micron. The size, density and shape of the particles determine their residence time in the atmosphere. Typically particles with a radius greater than 100 μm have low residence times (from minutes to hours), while particles smaller than 10 μm and gaseous aerosols can stay in the atmosphere for weeks and travel for thousands of kilometers.

Impact of volcanic eruptions

- **Health** [Horwell et al., 2006]
- ✓ Ballistic missiles
- ✓ Lava flow and pyroclastic flows [gas mixtures, ash, pumice and rocks pushed by gravity downstream along the flanks of the volcano]
- ✓ Lahar [mudflow composed of pyroclastic material and water flowing down the slopes of a volcano]
- ✓ Ash/Gas emissions
- ✓ Seismic activity, tsunamis

	H	Etna volatile fluxes (197	5-1995)		
Specie	Etna mean (Gg/yr)	Global Volcanism (Gg/yr)	Etna/Global Volcanism	Etna/Global Natural	
H ₂ O	5.0·10 ⁵	5.0·106	10%		
CO ₂	1.3-104	(8-20)·10 ⁴	(7-16)%	0.018%	
SO ₂	2.0·10 ³	1.3·10 ⁴	11%	0.68%	
HCl	4.8·10 ²	(4-110)·10 ²	9%	0.16%	
HF	59	60-6000	8%	0.70%	
Br	2	77	2.6%	0.19%	
Zn	4.9	9.6	51%	0.9%	
Cu	0.57	9.4	5.9%	≤ 1.0%	
Mn	0.23	42	0.6%	0.013%	
Pb	0.15	3.3	4.5%	1.25%	
Mo	0.08	0.08 0.4 20%		2.8%	
As	0.04	3.8	1.1%	0.33%	
Se	0.018	0.95	1.9%	2%	
Ni 0.036		3.8	1.0%	0.12%	
Hg	0.027	1.0	2.7%	1.1%	
Cd	0.050	0.82	6.1%	3.9%	
Sr	0.007	15	0.05%	0.016%	
Sb	0.004	0.71	0.56%	0.017%	

[P. Allard, private Communication]

Environment

[Thordasson et al., 2003; Grainger et al., 2003; Guo et al., 2004]

Volcanic eruptions can have a high impact on the environment by altering the climate:

- ✓ Large eruptions lead gas and ash into the atmosphere which may remain for several years. The main effect is to reflect solar radiation causing a net cooling effect
- ✓ Aerosols promote ozone destruction. In the troposphere, these aerosols may act as nuclei for the cirrus cloudy formation, further interfering with the irradiation of the Earth
- ✓ The majority of the SO₂ is converted into sulfuric acid which condenses into fine particulate. These sulfur aerosols reflect solar radiation, cooling the troposphere, and absorb heat from the earth, warming the stratosphere
- ✓ CO₂ is a greenhouse gas and contributes to global warming

Courtesy from F. Prata

Mass of aerosol (Tg)

HCI and HF are soluble and fall to the ground as acid rain

Why study volcanic emissions? SO₂ and climate

SO₃ + H₃O + V - > H₃SO₄ aerosol is main cause for warming in stratosphere and cooling at the surface

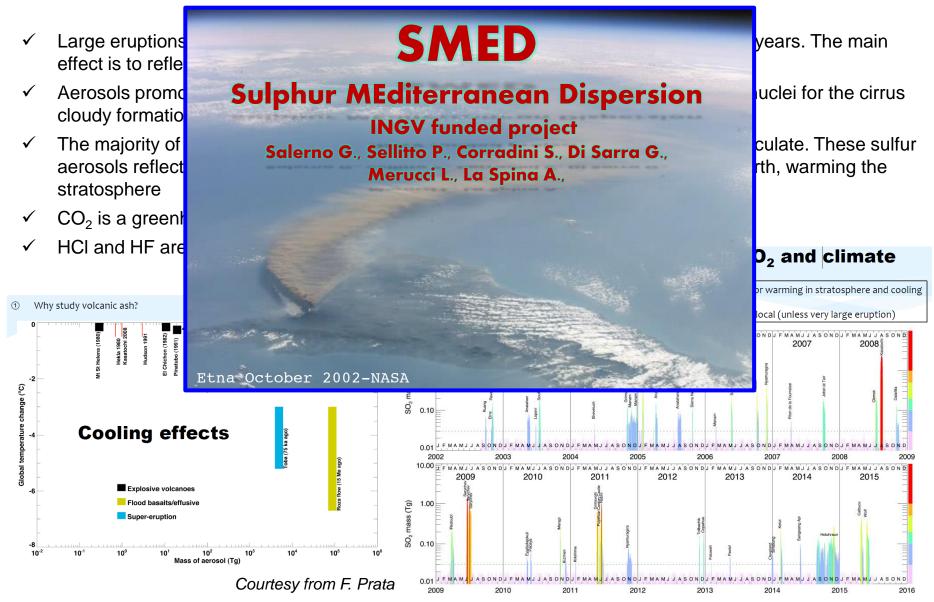
"Dust" can cause cooling but most likely only local (unless very large eruption)

To be a surface of the surface of t

Environment

[Thordasson et al., 2003; Grainger et al., 2003; Guo et al., 2004]

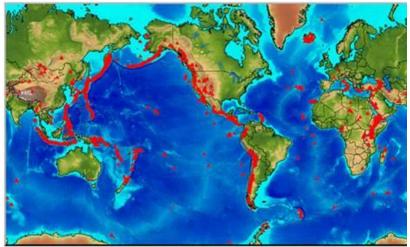
Volcanic eruptions can have a high impact on the environment by altering the climate:



- Air flight [Casadevall, 1994; Prata, 2009; Zehner, 2012; Prata and Prata, 2015]
 - Loss of engine power
 - Engine shutdown turbine
 - Abrasion of turbine blades
- Abrasion of the windows and windscreen
- Abrasion of the fuselage







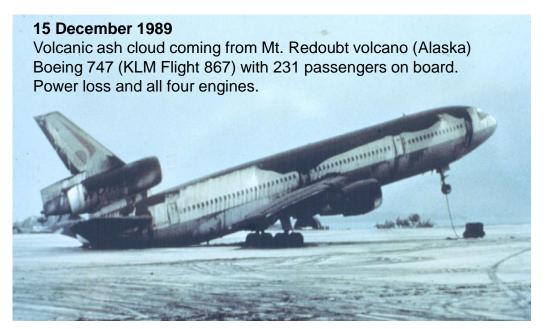
Total incidents reported

Encounter Severity Index

129

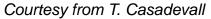
Sev	erity Class	Numbe
engine failure or other damage leading to crash	5	0
temporary engine failure	4 (engine failure)	9
engine vibration and "surging"; engine damage	3	17
heavy cabin dust; exterior abrasion; deposition of ash in engines	2	53
light dust in cabin; engine temp. fluctuate but remain "normal"	1	12
sulfur odor and anomalous haze	0	23
Courtesy from T. Casadevall Lack	king data	15

Courtesy from T. Casadevall



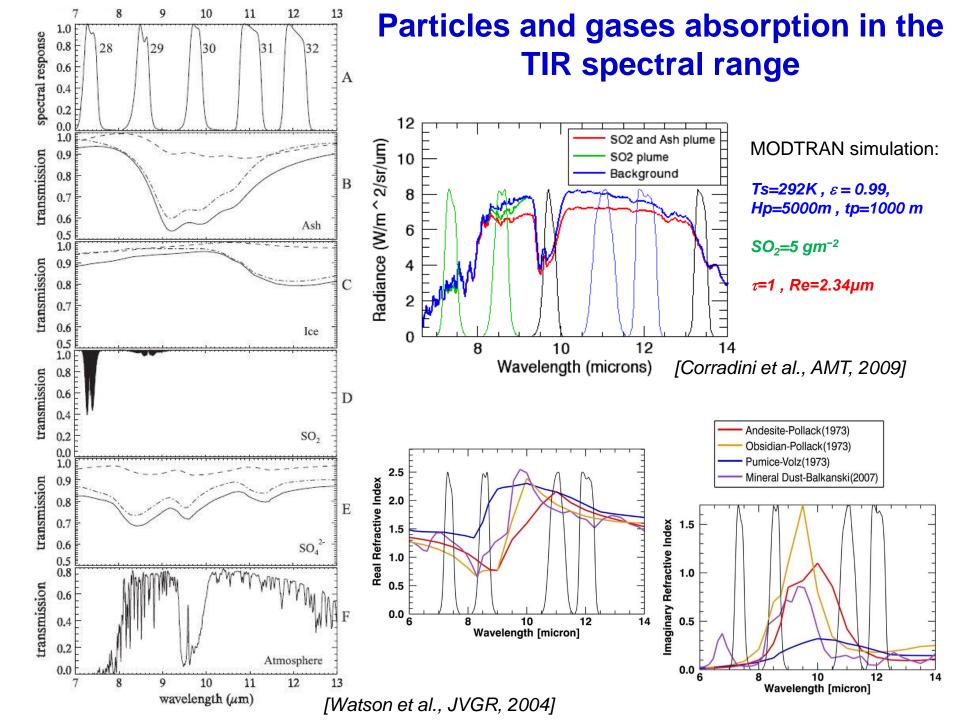




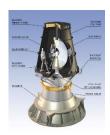




Courtesy from T. Casadevall



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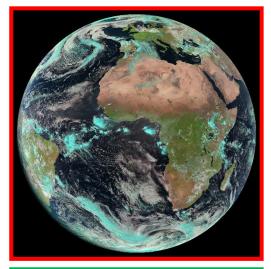
Spinning Enhanced Visible and Infrared Imager (SEVIRI)

	Temporal Resolution (min)	Spatial Resolution (km)	Central Wavelength (µm)	Channel No
	15	3	0.6	1
	15	3	0.8	2
	15	3	1.6	3
	15	3	3.9	4
	15	3	6.2	5
60	15	3	7.3	6
SO ₂	15	3	8.7	7
	15	3	9.7	8
Ash	15	3	10.8	9
ASII	15	3	12	10
	15	3	13.4	11
	15	1	0.8	12 (HRV)

Current Meteosat Satellites

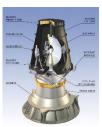
SATELLITE	LIFETIME	POSITION	SERVICES	
MSG-4	15/07/2015 – tbc	In orbit storage	n/a	
Meteosat-10 (MSG)	05/07/2012 – Nominal fuel lifetime is until 2022	0°/36,000 km	0º SEVIRI Image Data. Real-time Imagery. Data Collection Service	
Meteosat-9	22/12/2005 – Fuel lifetime is expected to be extended until 2021	9.5° E/36,000	Rapid Scan Service from 9 April 2013.	
(MSG)		km	Real-time Imagery	
Meteosat-8	28/08/2002 – Fuel lifetime is expected to be extended until 2019	3.5° E/36,000	Backup service for 0°, plus RSS from 9	
(MSG)		km	April 2013	
Meteosat-7	02/09/1997 (IODC since 01/11/2006) –	57° E/36,000	Indian Ocean Coverage. Real-time	
(MFG)	2016	km	Imagery. Date Collection Service	

http://www.eumetsat.int/website/home/Satellites/index.html

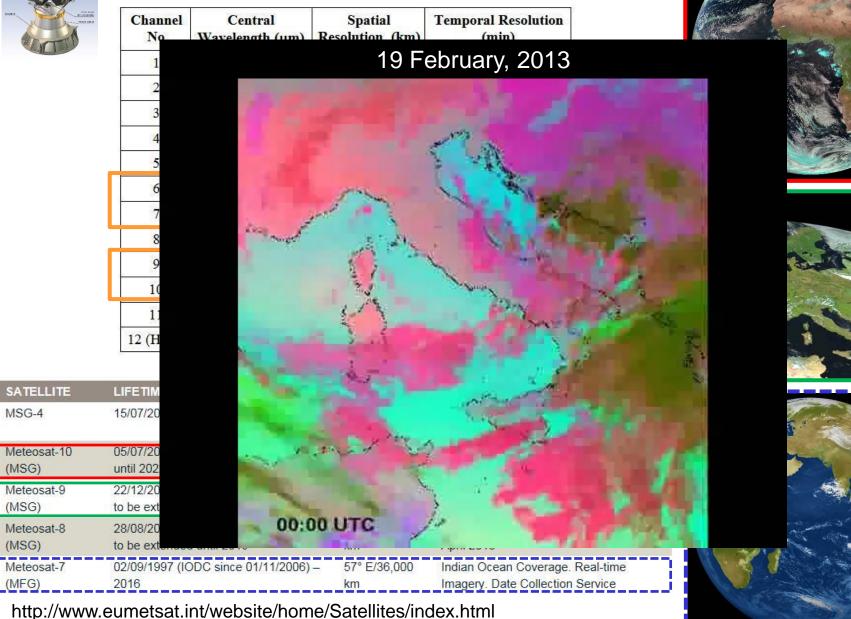








Spinning Enhanced Visible and Infrared Imager (SEVIRI)



Primary Use	Band	Bandwidth ¹	Spectral Radiance ²	Required SNR ³		MC	DIS		
Land/Cloud/Aerosols	1	620 - 670	21.8	128	1	Litter Herrich and a management			
Boundaries	2	841 - 876	24.7	201	http://modis.gsfc.nasa.gov/				
_and/Cloud/Aerosols	3	459 - 479	35.3	la a	-				
Properties	4	545 - 565	29.0	29.0 Primary Use	Band	Bandwidth ¹	Spectral Radiance ²	Required NE[delta]T(K) ⁴	
	5	1230 - 1250	5.4	Surface/Cloud	20	3.660 - 3.840	0.45(300K)	0.05	
	6	1628 - 1652	7.3	Temperature	21	3.929 - 3.989	2.38(335K)	2.00	
	7	2105 - 2155	1.0		22	3.929 - 3.989	0.67(300K)	0.07	
Ocean Color/	8	405 - 420	44.9		23	4.020 - 4.080	0.79(300K)	0.07	
Phytoplankton/ Biogeochemistry	9	438 - 448	41.9	Atmospheric Temperature	24	4.433 - 4.498	0.17(250K)	0.25	
biogeochemistry	10	483 - 493	32.1		25	4.482 - 4.549	0.59(275K)	0.25	
	11	526 - 536	27.9	Cirrus Clouds Water Vapor	26	1.360 - 1.390	6.00	150(SNR)	
	12	546 - 556	21.0		27	6 535 - 6 895	1.16(240K)	0.25	
	13	662 - 672	9.5		28	7.175 - 7.475	2.18(250K)	0.25	
	14	673 - 683	8.7	Cloud Properties	29	8.400 - 8.700	9.58(300K)	0.05	
	15	743 - 753	10.2	Ozone	30	9.580 - 9.880	3.69(250K)	0.25	
	16	862 - 877	6.2	Surface/Cloud	31	10.780 - 11.280	9.55(300K)	0.05 Ash	
Atmospheric Water Vapor	17	890 - 920	10.0	Temperature	32	11.770 - 12.270	8.94(300K)	0.05	
	18	931 - 941	3.6	Cloud Top Altitude	33	13.185 - 13.485	4.52(260K)	0.25	
	19	915 - 965	15.0		34	13.485 - 13.785	3.76(250K)	0.25	
					35	13.785 - 14.085	3.11(240K)	0.25	

MODIS on board the NASA-Terra & NASA-Aqua satellites

Launch Date: December 18,1999

Equatorial crossing: 10:30 a.m. descending node

Launch Date: May 4, 2002

Equatorial crossing: 1:30 p.m. ascending node

	36	14.085 - 14.385	2.08(220K)	0.35
Altitude		705 km		
Inclination	!	99°		
Period	!	99 minutes (16	orbits per	day)
Orbit		Sun synchronoι	us, near po	lar orbit
Mapping		1-2 days		
Swath Wid	th :	230 km		
Spatial Res	. :	1 km (TIR-nadir)	

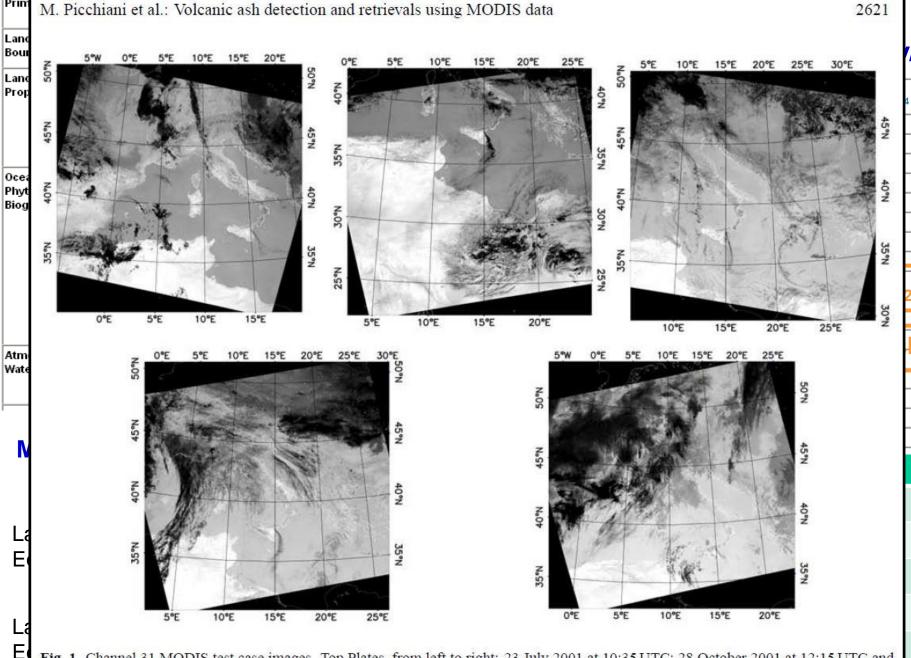
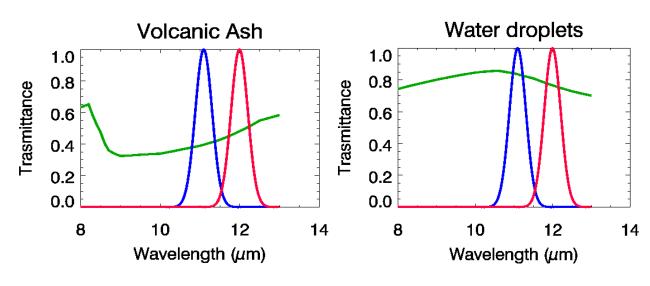


Fig. 1. Channel 31 MODIS test case images. Top Plates, from left to right: 23 July 2001 at 10:35 UTC; 28 October 2001 at 12:15 UTC and 29 October 2002 at 12:05 UTC. Bottom Plates, from left to right: 30 October 2002 at 9:45 UTC and 24 November 2006 at 12:20 UTC.

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Ash Detection, the BTD approach [Prata, 1989]

The **BTD** (Brightness Temperature Difference) technique is based on the different absorption of ash and ice/water droplets in the channels centered at 11 and 12 µm



BTD =
$$T_b(11 \mu m) - T_b(12 \mu m)$$

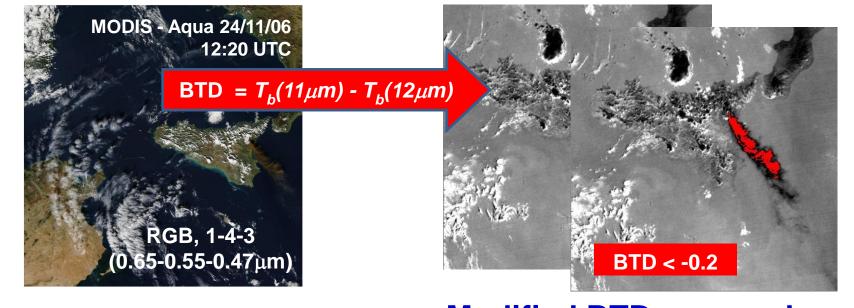
$$T_b(11 \mu m) - T_b(12 \mu m) < \alpha (<0)$$
 $T_b(11 \mu m) - T_b(12 \mu m) > \beta (>0)$

$$T_b(11 \mu m) - T_b(12 \mu m) > \beta (>0)$$

Advantages:

Very simple application

- Drawbacks (where the procedure fails) [Prata et al., 2002]:
- Over land surfaces during night;
- In the presence of high content of water vapour;
- On soils containing high amounts of quartz (deserts);
- On very cold surfaces (temperatures below 220 K);
- On surfaces covered by ice.

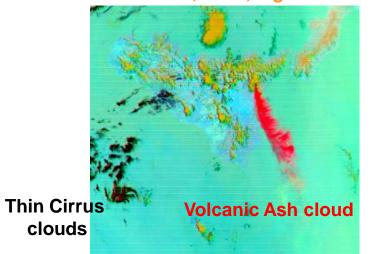


Red-Green-Blue (RGB) composite image Modified BTD approaches
[Millington et al., 2012]

R: $T_b(12 \mu m) - T_b(11 \mu m)$ G: $T_b(11 \mu m) - T_b(8.7 \mu m)$

B: $T_b(10.8 \mu m)$

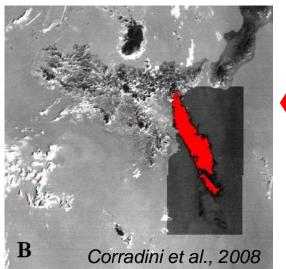
Cold, thick, high-level clouds



Atmospheric water vapour correction

[Prata et al., 2001; Yu et al., 2002;

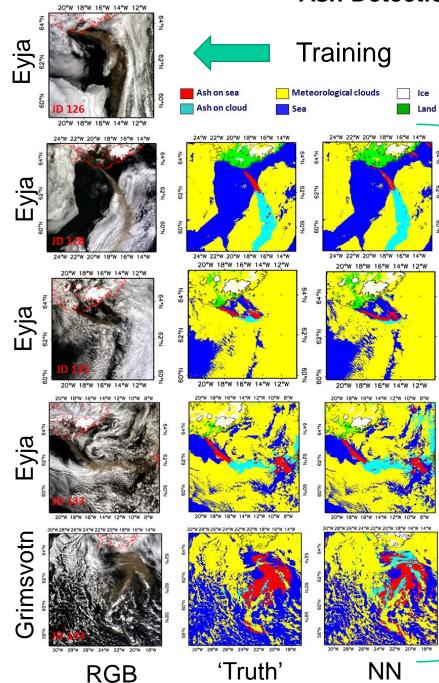
Corradini et al., 2008]





3069 1292 pixels

Ash Detection by using Neural Network approach [Picchiani et al. 2014]



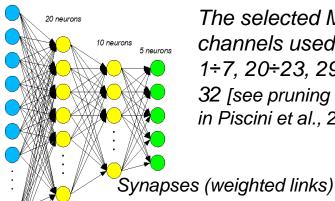
The NN has been trained using:

BTD approach

30 neurons

- MODIS L1 Land/Sea Mask
- MODIS L2 Cloud Products
- Threshold on MODIS band 7 for ice surfaces

Multilayer Perceptron (MLP) NN [Bishop, 1995]

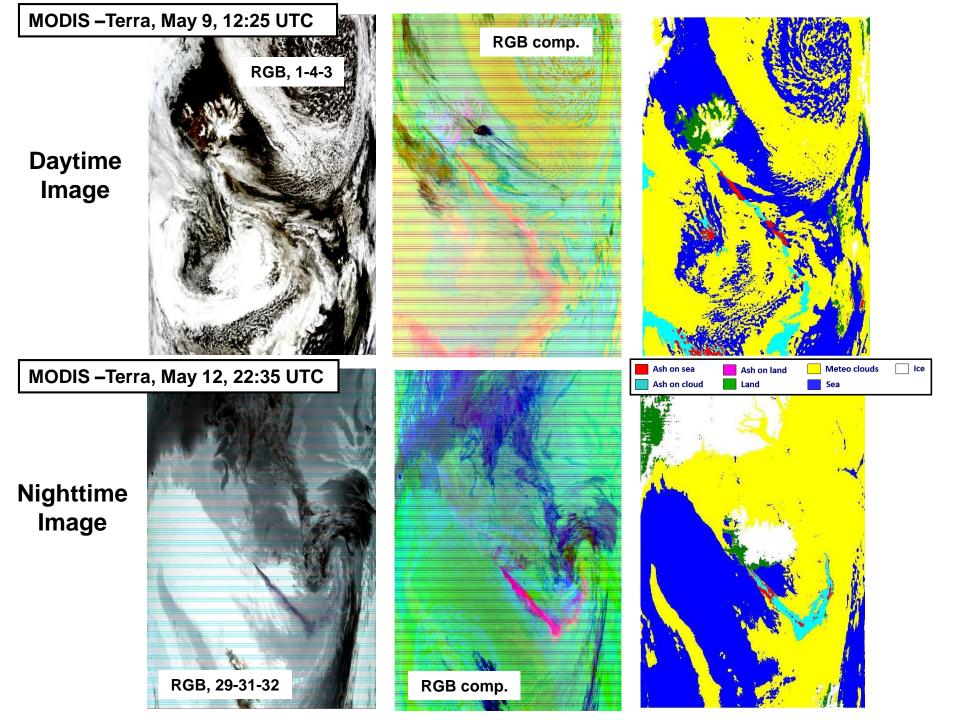


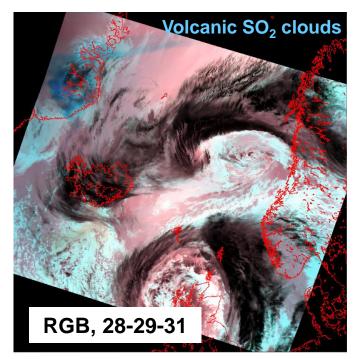
The selected MODIS channels used are: 1÷7, 20÷23, 29, 31 and 32 [see pruning analysis in Piscini et al., 2014].

Validation

Improvements:

- application of a smoothing filter to avoid stripe noise
- >training of a new neural networks for the nighttime images
- >new class: ash on land



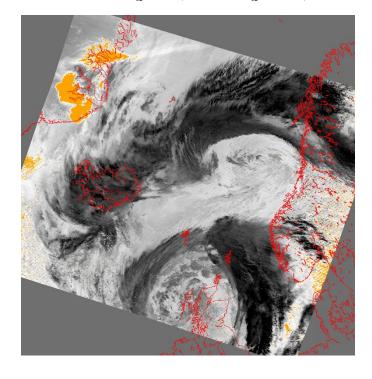


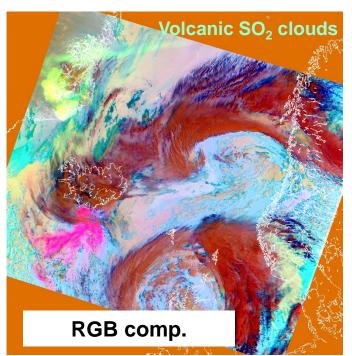
SO₂ detection

Grimsvotn (Iceland) 2011 eruption

MODIS-Terra, 23 May, 13:45 UTC





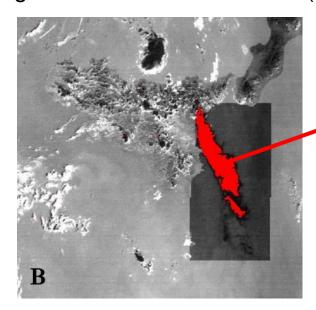


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The Look Up Tabel (LUT) approach: Ash retrievals

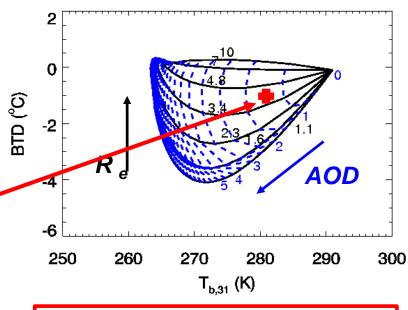
[Prata, 1989; Wen and Rose, 1994; Yu et al., 2002; Corradini et al., 2008; 2009; 2010]

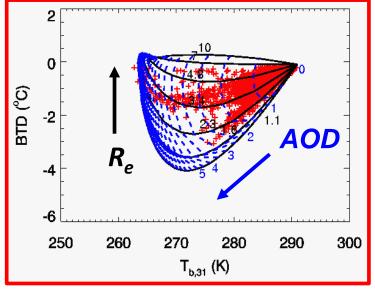
✓ The ash retrievals are based on computing the simulated inverted arches curves
"BTD - T_b(11µm)"
varying AOD and effective radius (R_e)



$$M_{(m,n)} = \frac{4}{3} S_{(m,n)} \rho \frac{R_{e(m,n)} AOD_{(m,n),\lambda}}{Q_{ext,\lambda}(R_e)}$$

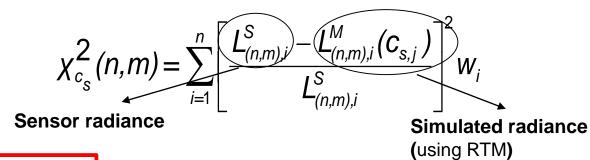
✓ The simulated inverted arches curves are computed using Radiative Transfer Models (RTM)

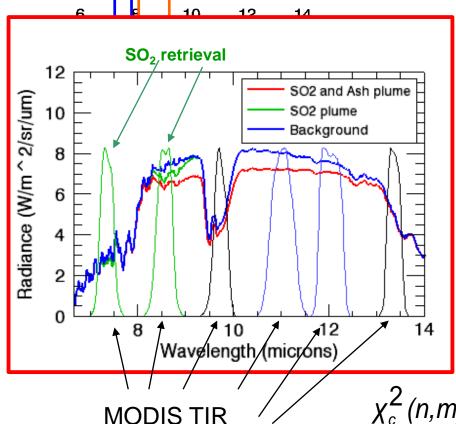




SO₂ retrieval

χ2 Procedure [Realmuto et al., JGR, 1994; Teggi et al., JGR, 1999]





response functions

SO,

8.6 μm

7.3 μm

1.0

0.8

0.6

0.4

0.2

0.0

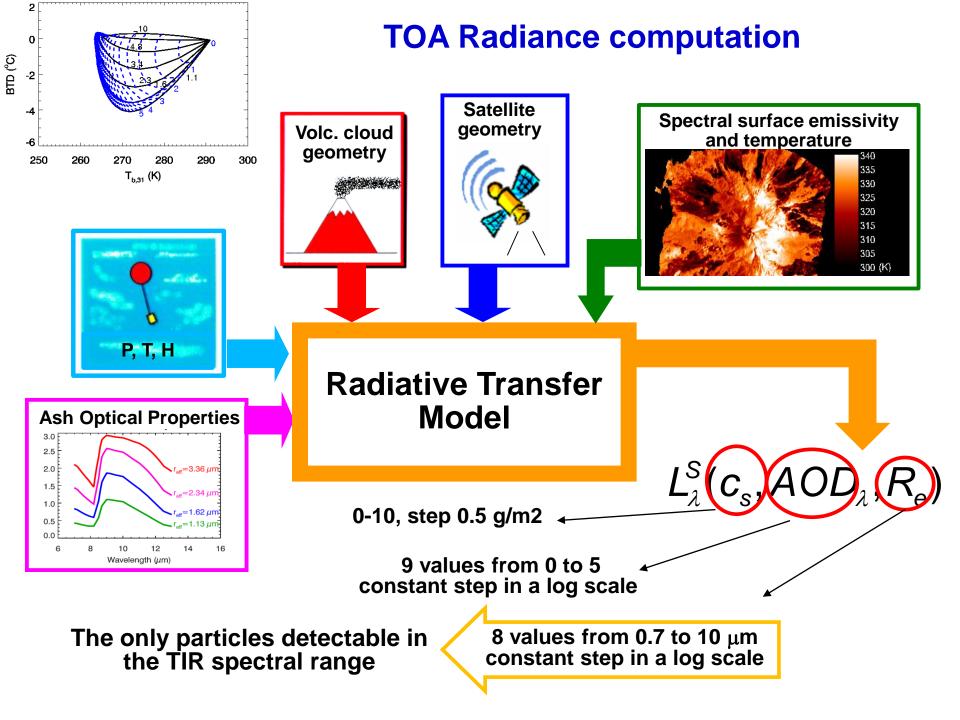
Trasmittance

Ash effect on SO₂ retrieval

- During an eruption generally ash and gases are emitted simultaneously
- ➤ The plume ash particles reduce the top of atmosphere radiance in the entire TIR spectral range, including the channels used for the SO₂ retrieval
- ➤ The net effect is a significant SO₂ overestimation

[Corradini et al., AMT, 2009; Kearney and Watson, JGR, 2009]

$$\chi_{c_s}^2(n,m) = \sum_{i=1}^n \left[\frac{L_{(n,m),i}^S - L_{(n,m),i}^M(c_{s,j}, AOD_{(n,m)}, R_{e(n,m)})}{L_{(n,m),i}^S} \right]^2 w_i$$



Etna (Italy) 2002 Kasatochi (Alaska) 2008 Eyjafjalla (Iceland) 2010 MODIS-Aqua 28 October 2002, 12:00 UTC MODIS-Aqua 9 August 2008, 00:50 UTC MODIS-Terra 19 April 2010, 14:55 UTC A00 (11µm) A00 (11µm) 8,0 Ash Mass (:/km²) Ash Mass (1/km²) Ash Mass (1/km²) [Corradini et al., 2010] [SAVAA project, http://savaa.nilu.no/] [Corradini et al., 2011]

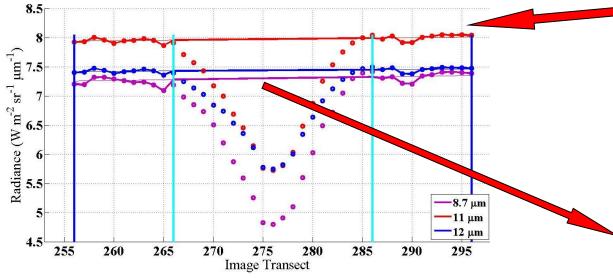
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The Volcanic Plume Removal (VPR) procedure

[Pugnaghi et al., AMT, 2013; Guerrieri et al., JVGR, 2015]

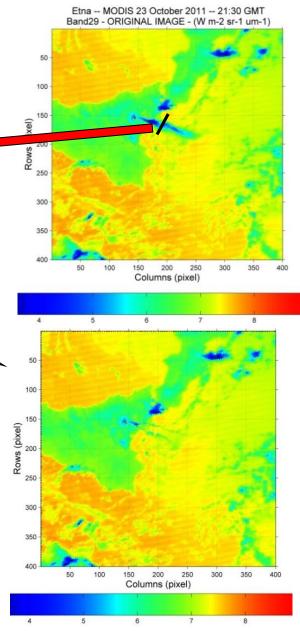
The volcanic cloud signal in the TIR spectral range can be seen as a region characterized by a dip in the radiance



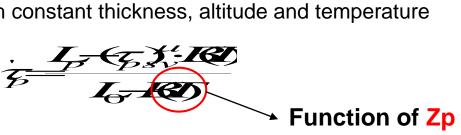


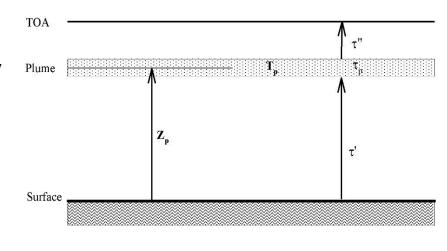
L0 = modified radiance without plume

Step 1) The VPR procedure starts with the removal of the plume from the image obtained by substituting the radiance values in the plume region with the interpolated values found in region surrounding the plume.

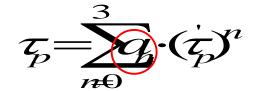


Step 2) The new image and the original data allow the computation of the **plume transmittance** (τ) in the TIR bands centred at 11 (τ'_{II}) and 12 (τ'_{I2}) µm by applying a simplified model consisting of a plume with constant thickness, altitude and temperature





Step 3) The plume transmittances are then refined with a **polynomial relationship** obtained by means of MODTRAN simulations **adapted for geographical region, ash type and atmospheric profiles**



Step 4) By knowing τ_{11} and τ_{12} we can retrieve the ash **AOD** and the **Re** and then the **Ash mass** with the Wen and Rose (1994) simplified formula:

$$M = \frac{4}{3} S \rho \frac{R_e A O D_{\lambda}}{Q_{ext,\lambda}(R_e)}$$

Step 5) And finally, by exploiting the relationship between the τ_{II} and $\tau_{g,\gamma}$ (only ash) we can compute $\tau_{g,\gamma}$ of the **SO₂** only:

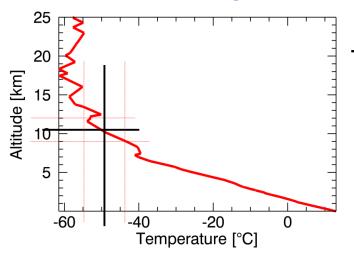
$$\tau_{p,SO2} = e^{-\mu \beta Cs}$$

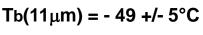
Once the coefficients have been computed for a specific area, volcano and instrument, the only input is the cloud altitude

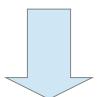
Volcanic cloud altitude computation

Dark pixels procedure:

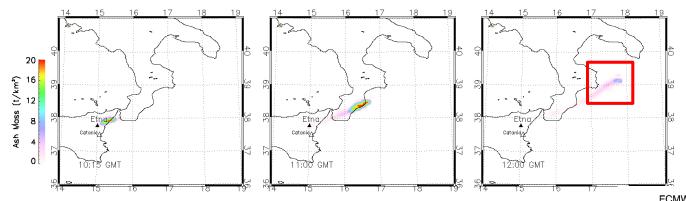
The volcanic cloud altitude is obtained by comparing the Tb at 11µm of the satellite image 'dark pixels', with the atmospheric temperature profile of the nearest WMO meteo station of the area of interest [Prata and Grant, 2001; Corradini et al., 2010; Scollo et al., 2014]



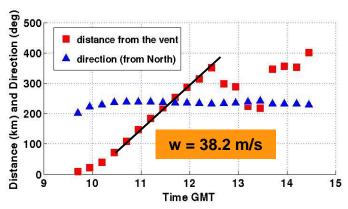


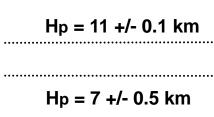


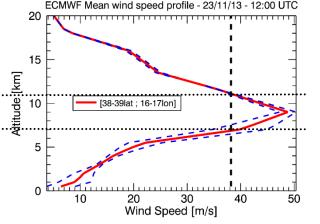
Hp = 10.1 + /- 1.3 km



Track of the volcanic cloud center of mass:







Summary ... and other techniques

Procedure	Characteristics
Dark Pixels	 Very simple application Not applicable for diluted cloud Tendency to underestimate the altitude
Center of Mass	Applicable only by using geostationary data
Cloud Shadow	 Applicable only during daytime Not applicable in case of diluted clouds
Satellite Parallax	Critic in case of diluted clouds
CO ₂ slicing	 For instruments with several channels inside the TIR CO₂ absorption band Critic in case of diluted clouds Tendency to underestimate the altitude

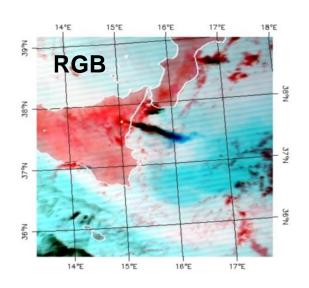
Test case event - October 23th 2011

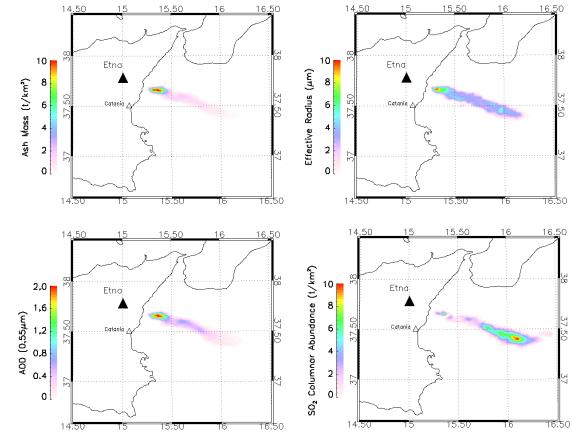
From INGV-Catania report: "on the evening of 23 October 2011, the 17th paroxysmal eruptive episode occurred at the New Southeast Crater (New SEC) of Etna. The cloud of gas and tephra was carried eastward by the wind, leading to ash and lapilli falls in a sector from Zafferana Etnea and Milo to Giarre and Riposto from the 19:40 to 21:00".

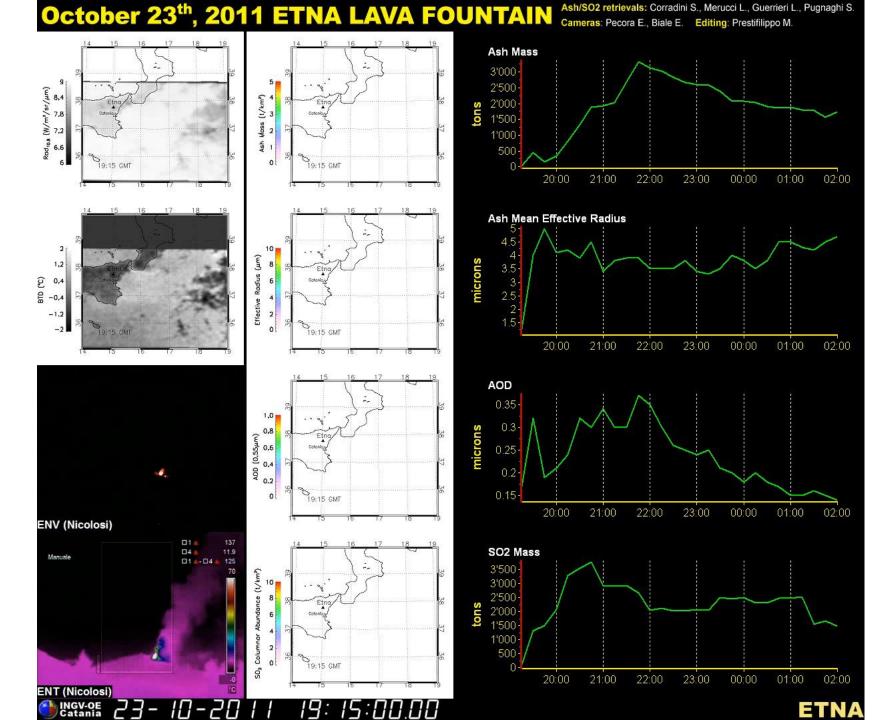


New south East crater, 20:40 GMT [B. Bencke]

MODIS (21:30 UTC)



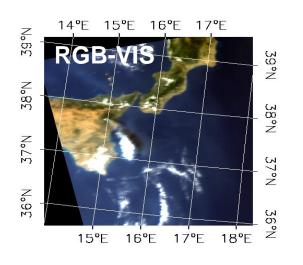


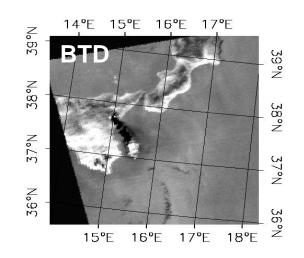


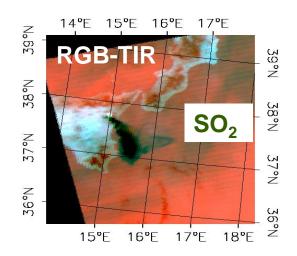
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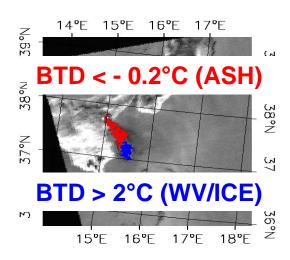
• The Etna 2011 lava fountains produced volcanic clouds containing ash, gases and water vapour/ice particles

Etna (Italy), 12 August 2011 MODIS-Aqua - 11:15 UTC









 In many cases the ash signal was completely hidden by the wv/ice particles

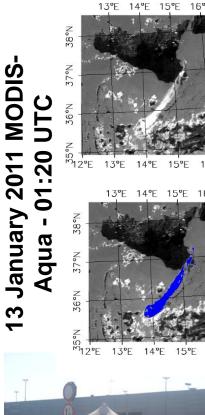
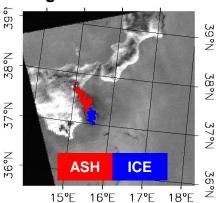


Photo by S. Scollo

1) ASH vs WV/Ice discrimination

(using the BTD approach)

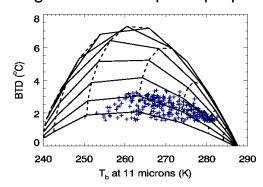
12 August 2011-11:15 UTC



3) WV/Ice retrievals

LUT:

computing the "BTD-Tb($11\mu m$)" curves using the WV/Ice optical properties



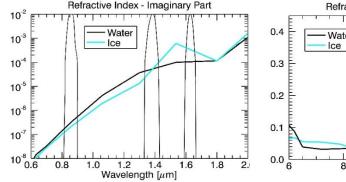
VPR:

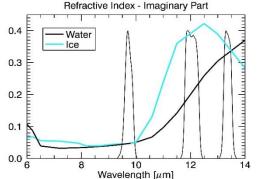
computing the coefficients by considering the WV/Ice optical properties

WV/Ice detection and retrieval

2) WV vs Ice discrimination

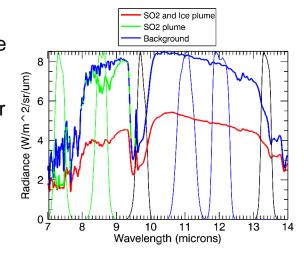
(using the VIS-SWIR channels during daytime)





WV/Ice effect on SO₂ retrieval

- As the plume ash particles, the WV/Ice particles reduce the top of atmosphere radiance in the entire TIR spectral range, including the channels used for the SO₂ retrieval
- The net effect is a significant SO₂ overestimation



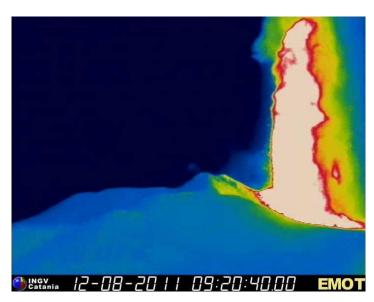
LUT:

$$\chi_{c_s}^2(n,m) = \sum_{i=1}^n \left[\frac{L_{(n,m),i}^S - L_{(n,m),i}^M(c_{s,j}, AOD_{(n,m)}^{ice}, R_{e,(n,m)}^{ice})}{L_{(n,m),i}^S} \right]^2 w_i$$

Test case event - August 12th 2011

From INGV-Catania reports [B. Bencke]: the lava fountain started at 8:30 UTC reaching an height of several hundred meters with ash and lapilli fall that reached the village of Zafferana Etnea. A plume of mixed ash and SO_2 was emitted at an altitude of about 7 km. The explosive activity decreased after 10:30 and from 11:00 faded with weak ash pulses lasted half an hour.

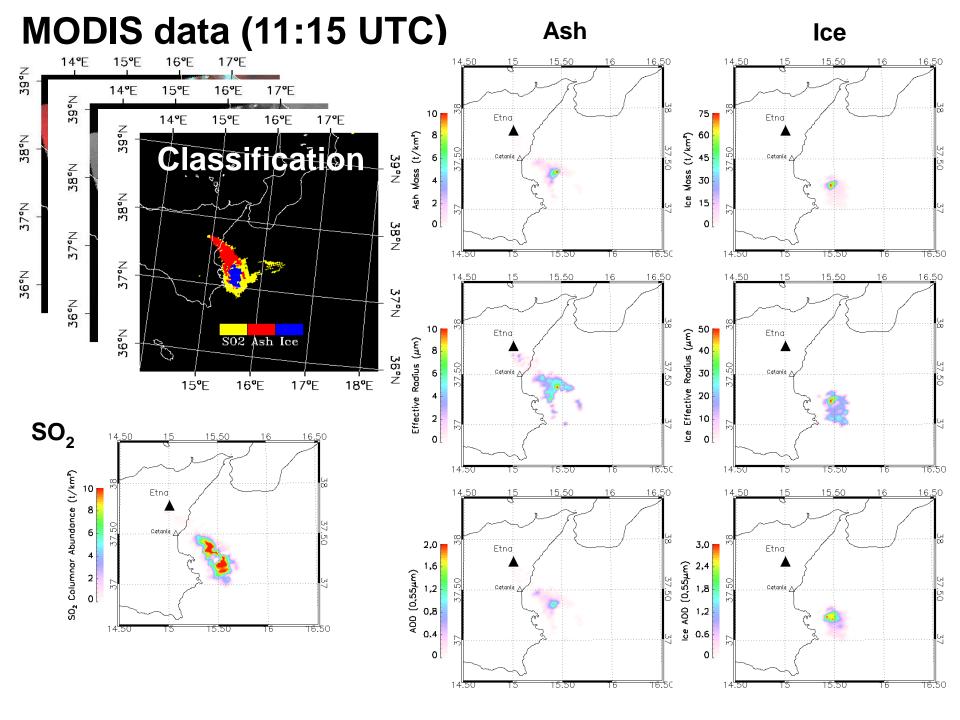
The satellite images emphasized the simultaneous presence of ash, SO₂ and Ice in the volcanic cloud.

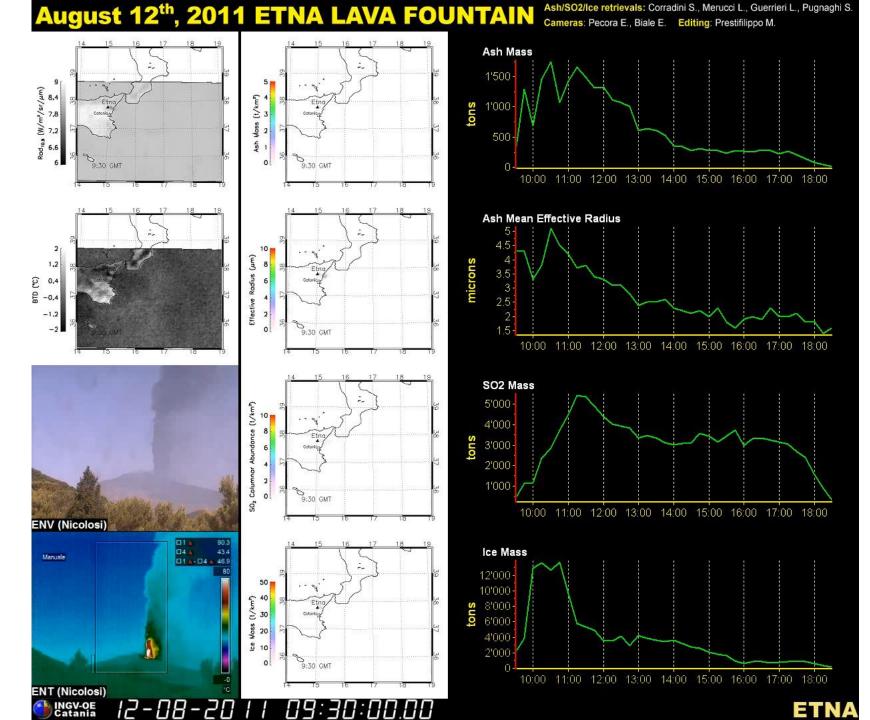


IR camera located on la Montagnola



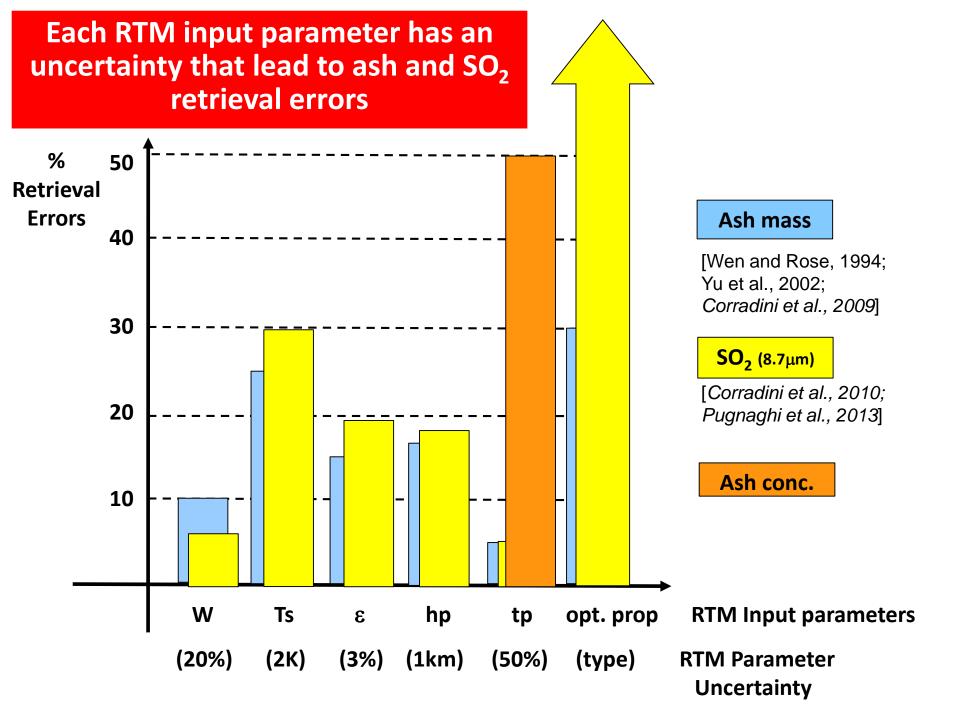
Courtesy from Elisabetta Ferrera, University of Catania





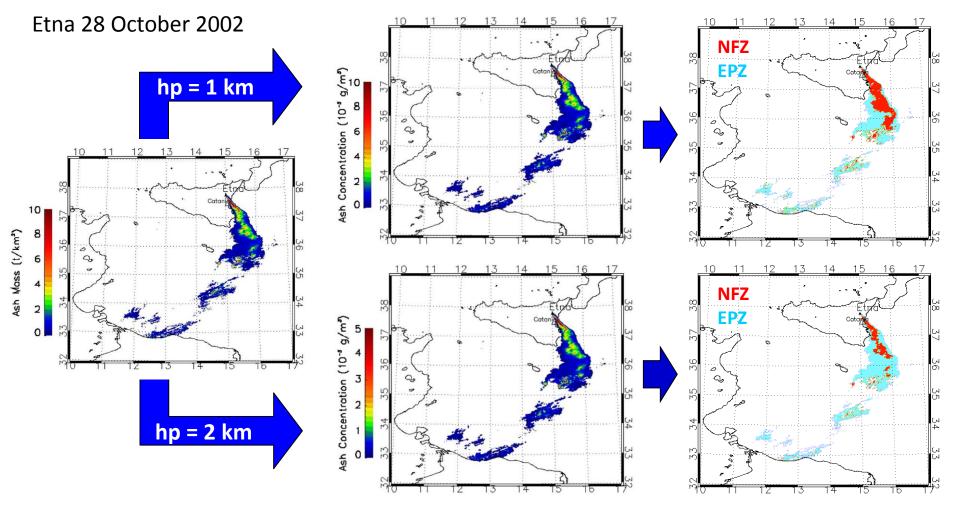
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- > Validation
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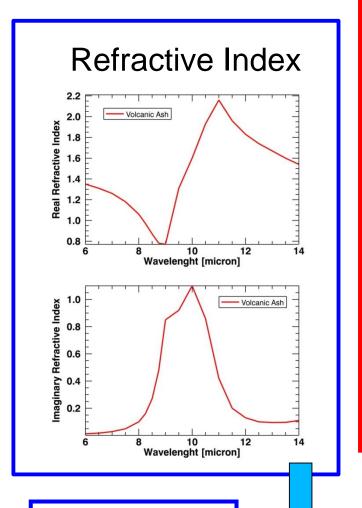
MODIS-Aqua,

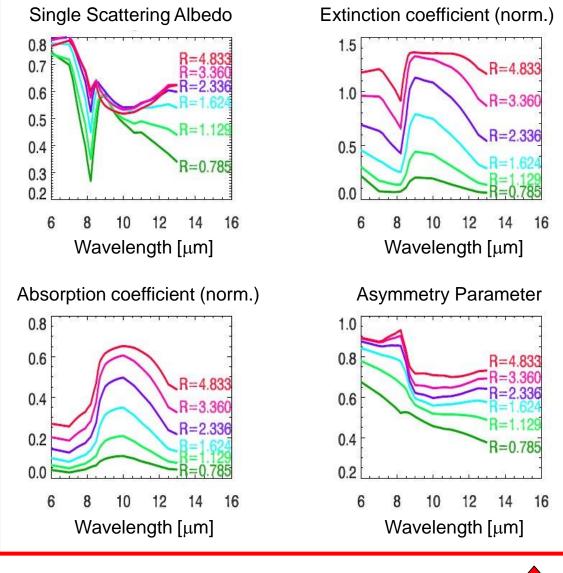
Ash cloud thickness



- No Fly Zone (NFZ): Any area where volcanic ash concentrations are predicted to be higher than $2 \times 10^{-3} \text{ gm}^{-3}$
- *Enhanced Procedures Zone (EPZ)*: Any area where volcanic ash concentrations are predicted to be between $2 \times 10^{-4} \text{ gm}^{-3}$ and $2 \times 10^{-3} \text{ gm}^{-3}$.

Ash Optical Properties Computation





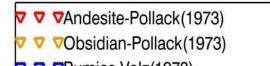
Size Distribution Mie Code (spherical approximation)

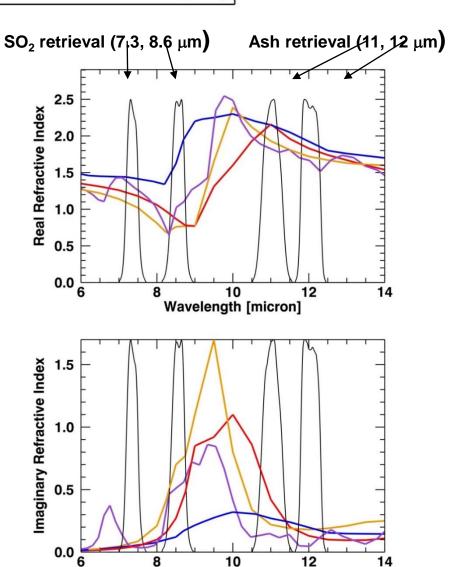
Andesite-Pollack(1973) Obsidian-Pollack(1973) Pumice-Volz(1973) Mineral Dust-Balkanski(2007)

Ash refractive index

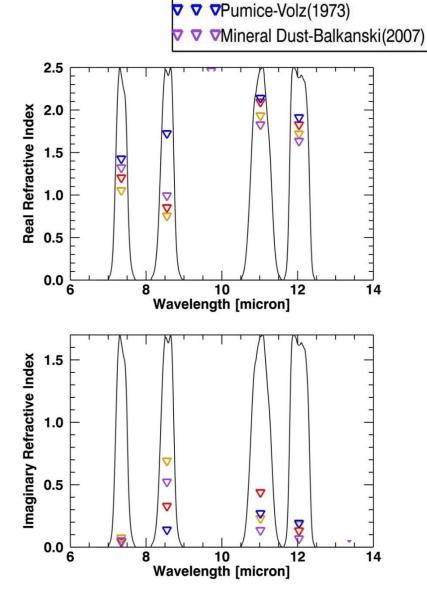
ARIA Database

http://www.atm.ox.ac.uk/project/RI/introduction.html





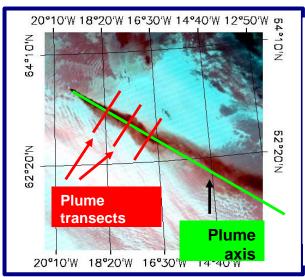
Wavelength [micron]

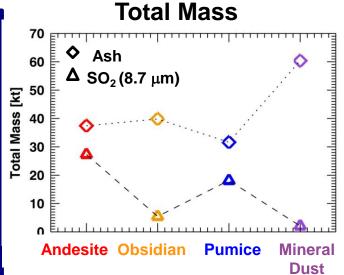


MODIS data collected on Eyja 2010 eruption

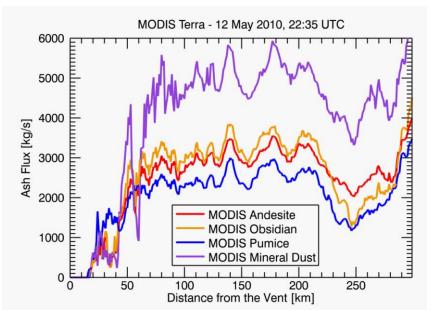
The 4 selected refractive indexes have been used to process the MODIS images collected the May 12, 2010 at 22:35 UTC

$$F_{t} = w_{s} I_{(n,m)} \sum_{t} m_{(n,m)}$$
$$\forall (n,m) \in t$$

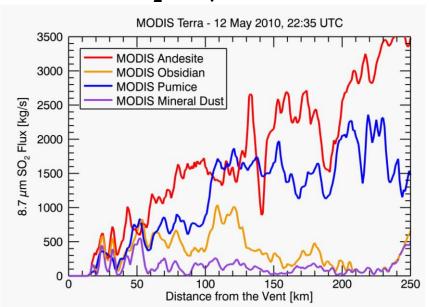




Ash Flux



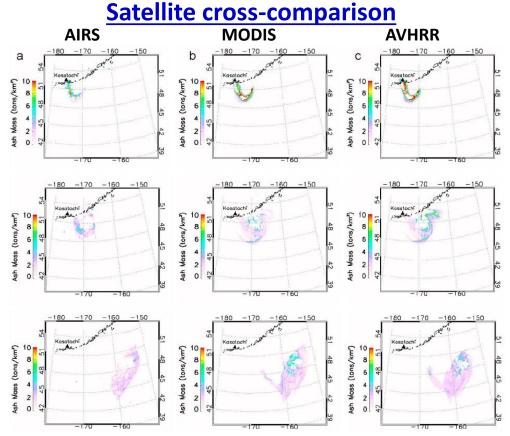
SO₂ 8.7 µm Flux



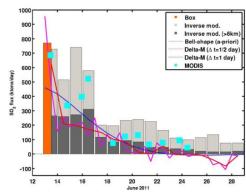
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Cross-Comparison and Validation



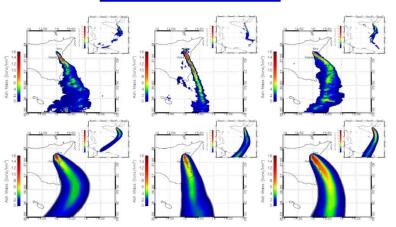
Kasatochi 2008 Eruption [Corradini et al., JGR, 2010]



Satellite SO₂ flux cross-comparison

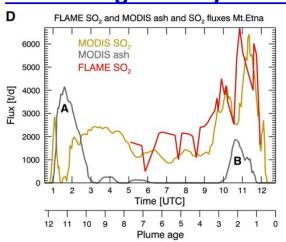
OMI, IASI, GOME2, MODIS, Nabro 2011 Eruption [Theys et al., ACP, 2013]

Satellite-Models



MODIS vs FALL3D Etna 2002 Eruption [Corradini et al., IEEE, 2011]

Satellite - ground systems



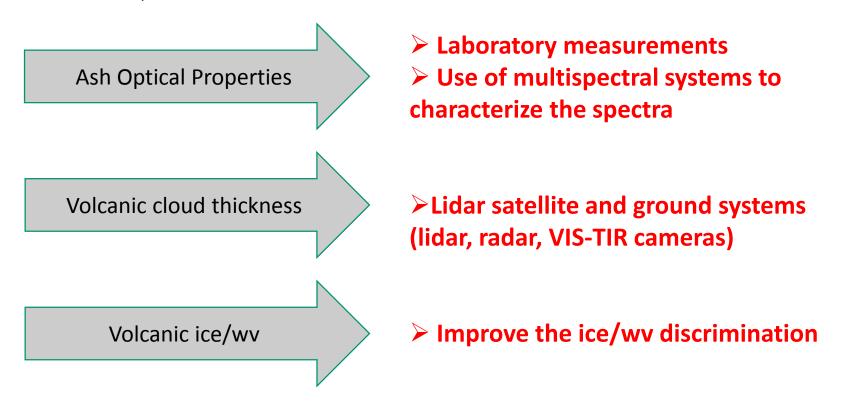
MODIS vs FLAME, Etna 2006 Eruption [Merucci et al., JVGR, 2011]

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✓ The use of the multispectral satellite sensors data collected from MODIS and SEVIRI have been proven to be efficient and the retrieval techniques effective

However, some criticalities still occur ...



✓ As one system cannot give a comprehensive description of the eruptive activity, exploit the capabilities to use different systems by integrating their measurements.

The APhoRISM FP7-EU Project

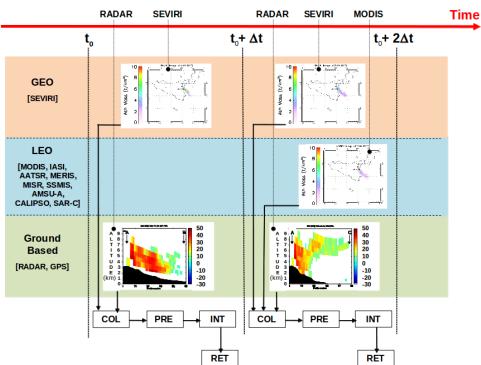
APhoRISM is a 3 years FP7-EU project started on December, 2013 that aims to *develop innovative products* to support the management and mitigation of the **volcanic** and the **seismic** crisis. Satellite and ground measurements will be managed in a novel manner to *provide new and improved products in terms of accuracy and quality of information*.



www.aphorism-project.eu/

The *Multi-platform volcanic Ash Cloud Estimation (MACE)* procedure will exploit the complementarity between geostationary, and polar satellite sensors and ground based measurements.

Mace is based on GEO-SEVIRI measurements. The basic idea is to continuously improve both the SEVIRI ash retrievals and the source characterization by integrating all the LEO and ground-based information from VIS to MW in clear and cloudy sky conditions.



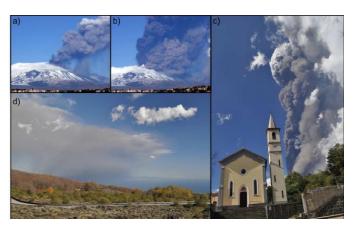
MACE temporal flow diagram

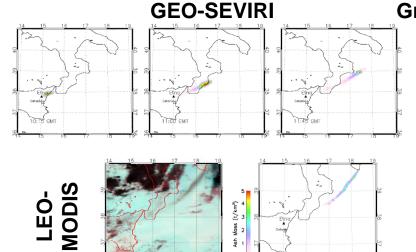
Test Cases

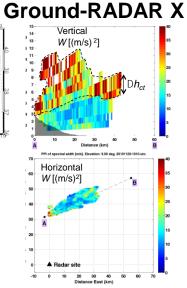
MACE ash products will be tested on 3 recent eruptions representative of different eruption styles in different clear or cloudy atmospheric conditions: Eyjafjallajokull (Iceland) 2010, Grimsvotn (Iceland) 2011 and Etna (Italy) 2011-2014.

23 November 2013 Etna lava fountain

This episode has been detected from:

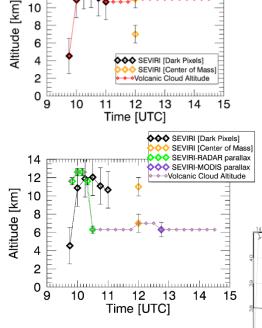




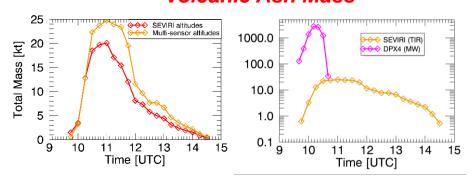


Volcanic cloud layers

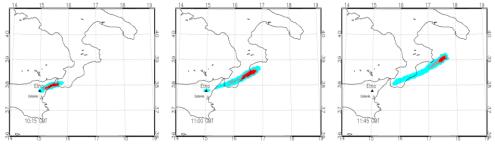
14 12



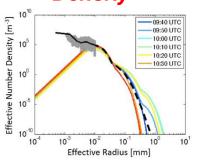
Volcanic Ash Mass



Volcanic Ash Concentration Classification



Effective Number Density



[Corradini et al., 2016]

